

Haline Convection
during Freezing of Seawater

Sönke Maus, Bergen 19.05.2006

- Rayleigh number
- Free convection
- Geophysical systems
- Sea ice

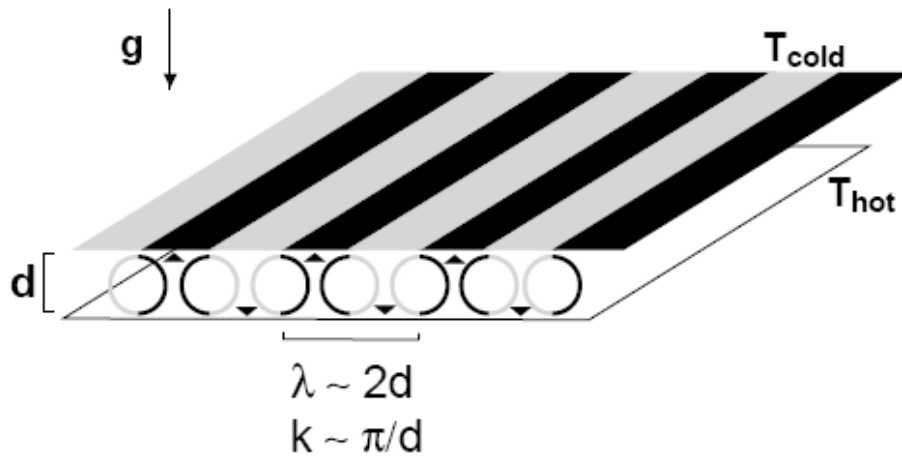
$$Ra = \frac{\beta \Delta S g H^3}{\nu D_s}$$

John William Strutt (1916)
Lord Rayleigh

$$Ra = \left(\frac{H}{H_*} \right)^3$$

$$H_* = \left(\frac{\nu D_s}{b} \right)^{1/3}$$

$$b = \beta \Delta S g$$



Silveston (1958)

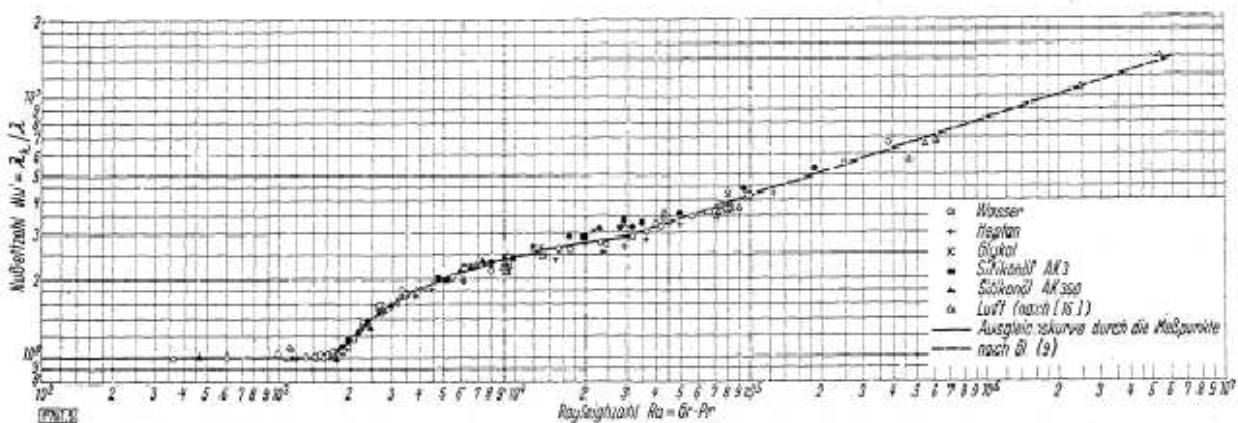


Bild 5. Abhängigkeit der Nußeltzahl Nu' von der Rayleighzahl Ra .
 λ und λ_k wahre bzw. scheinbare Wärmeleitzahl, Gr Grashof-, Pr Prandtlzahl

(rigid, rigid, $T = \text{const.}$, $T = \text{const.}$)

$$Ra_c \approx 1700$$

Chandrasekhar (1961)

”Hydrodynamic and Hydromagnetic Instability”

Rotation, Magnetic fields, Couette flow, Kelvin-Helmholtz

Superposed Fluids, Jets and Cylinders

Rayleigh, 1916 (free, free , T=const. , T=const.)

$$Ra_c = 657.511$$

Reid/ Harris, 1958 (rig, rig , T=const. , T=const.)

$$Ra_c = 1707.765$$

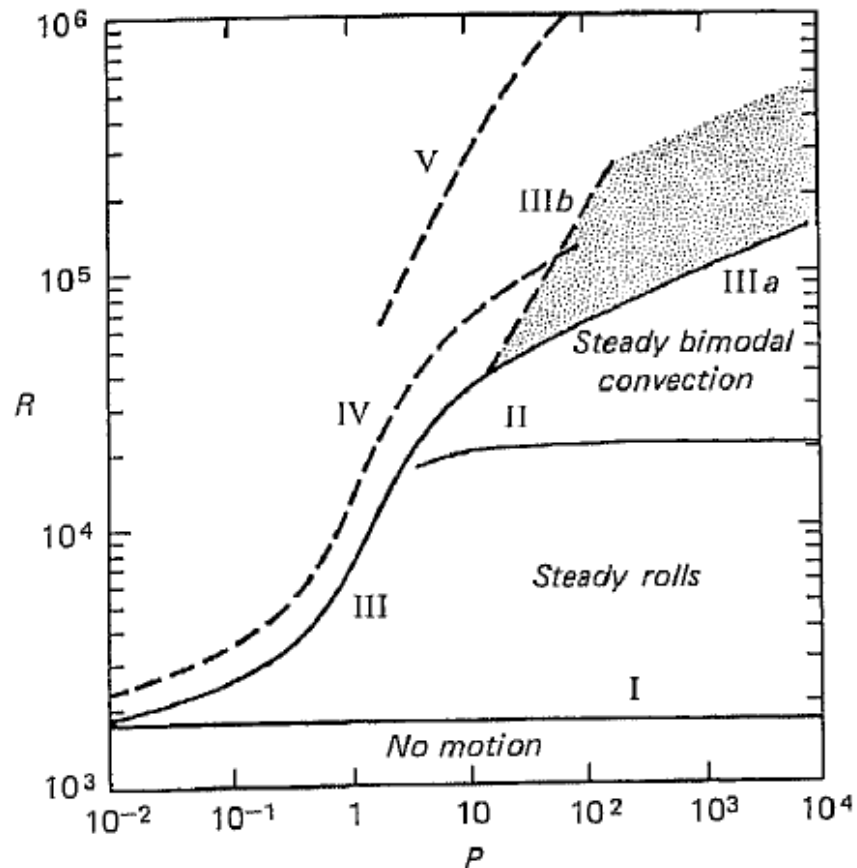
Nield, 1967 (free, free , Q=const. , Q=const.)

$$Ra_c = 120$$

$$4 + 3 + 2 + 1 = 10 \text{ combinations}$$

Transition to turbulence

Schmidt and Saunders (1938), Malkus (1954), Krishnamurti (1973)



Krishnamurti (1973)

Transition to oscillating flow depends on Pr

Free convection: $Ra \rightarrow \infty$

$$Nu \sim Ra^{1/3}$$

||: Saunders (1939), =: Weise (1940)

$$\underline{Ra \Rightarrow \infty \rightarrow Q \rightarrow const.}$$

Interfacial flux Q must become
independent of depth H

$$Nu = c_* Ra^{1/3} = c_* \frac{H}{H_*}$$

$$Q = Nu \frac{D_s \Delta S}{H}$$

$$\Rightarrow Q = c_* \frac{D_s \Delta S}{H_*}$$

$$H_* = \left(\frac{\nu D_s}{g \beta \Delta S} \right)^{1/3}$$

$$\Rightarrow Q = c_* \Delta S^{4/3} \left(\frac{D_s^2 g \beta}{\nu} \right)^{1/3}$$

Nonlinear flux law

$$Q = c_* \Delta S^{4/3} \left(\frac{D_s^2 g \beta}{\nu} \right)^{1/3}$$

How large is c_* ?

Intermittent convection

Howard's (1966) simple concept

$$c_* \approx Ra_c^{-1/3}$$

Howard: $Ra_c \approx 1000 \Rightarrow c_* \approx 0.1$

Generally, ad-hoc:

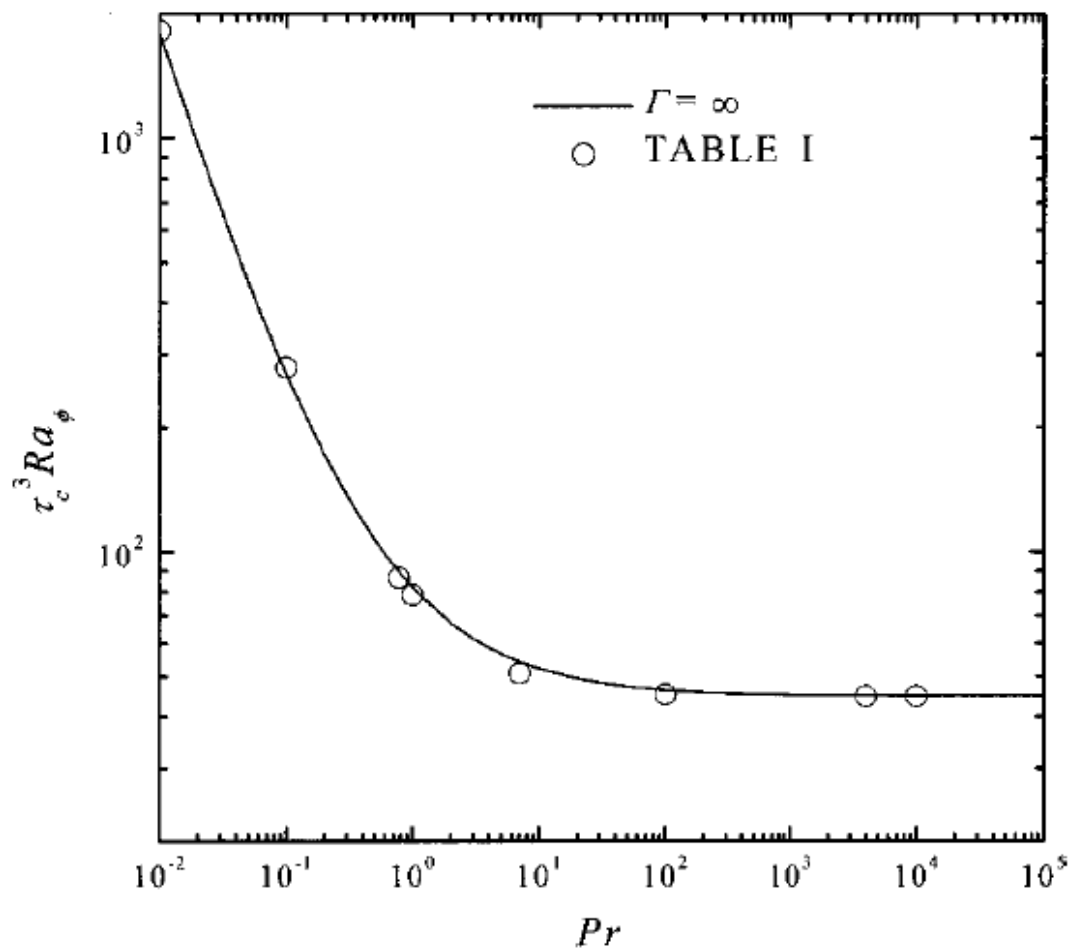
$$120 < Ra_c < 1708$$

\Rightarrow

$$0.08 < c_* < 0.28$$

Q and c_* depend on

- boundary conditions
- Prandtl Number



Yang and Choi (2002, Phys. Fluids)

See also: Foster (1968, Phys. Fluids),

Different Boundary Conditions

Free, rigid, mixed

Flux, double-diffusive

Geophysical Applications

Unstable Atmosphere

(Herring, Priestley, Deardorff)

Convection in Stars

(Spiegel)

Oceanic double diffusion

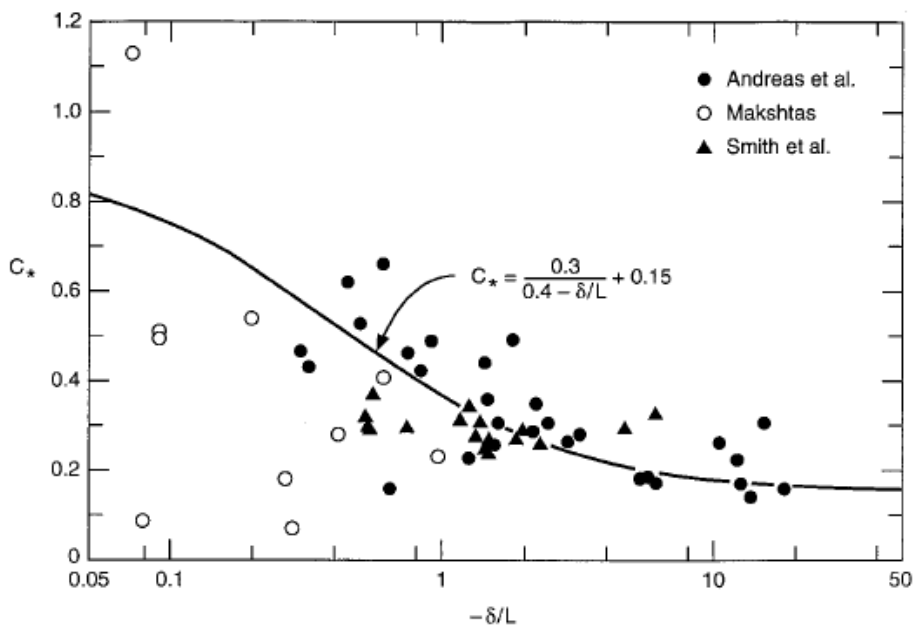
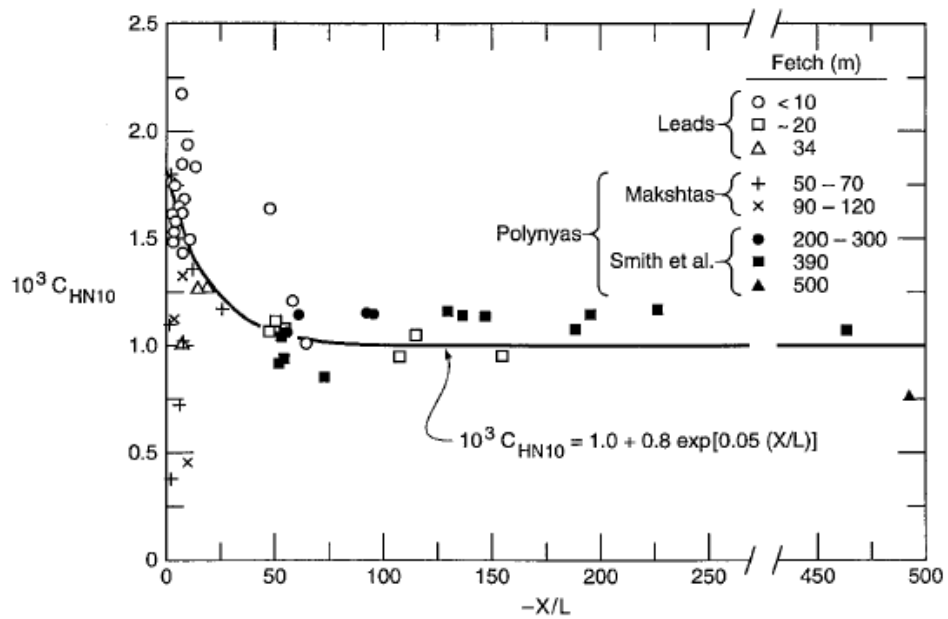
(Stommel, Stern)

Sea ice freezing

(Foster)

Arctic Leads

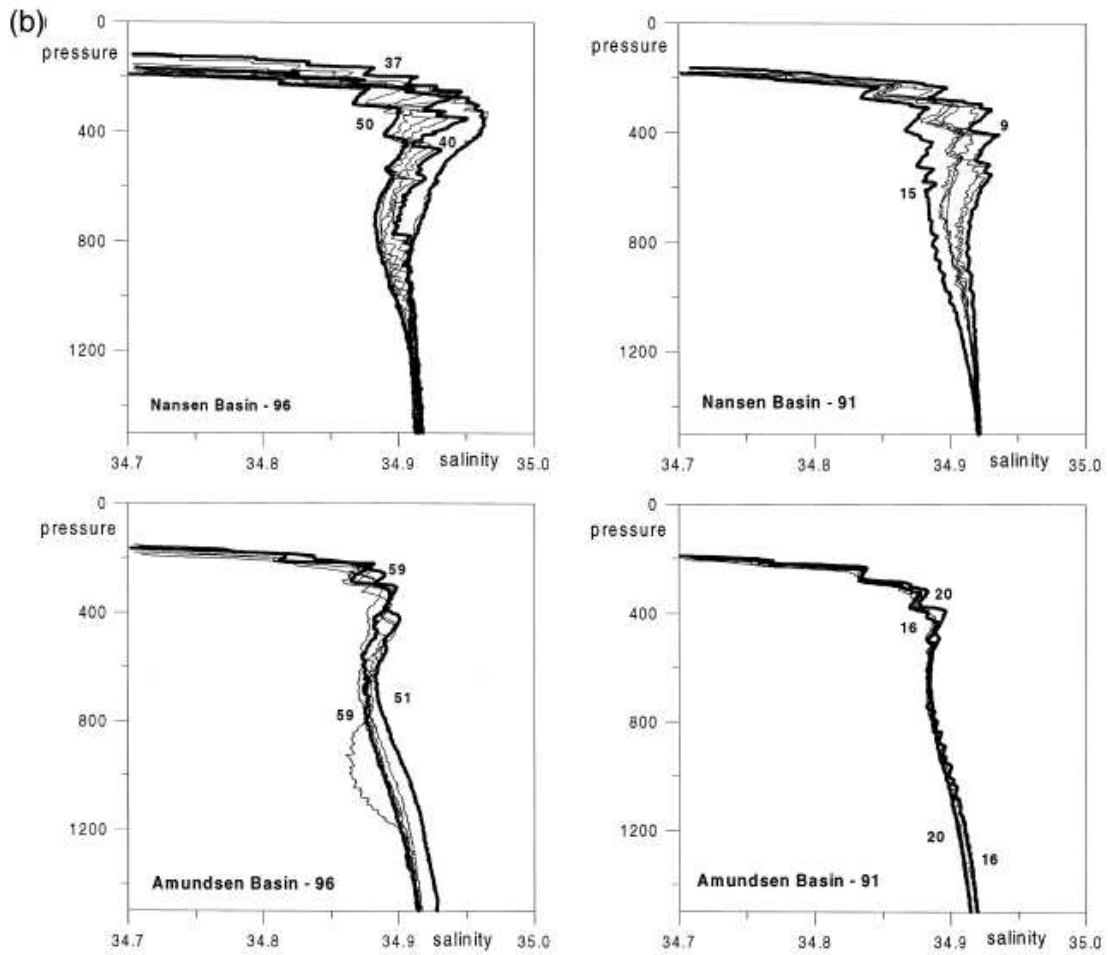
Austausch Coefficients



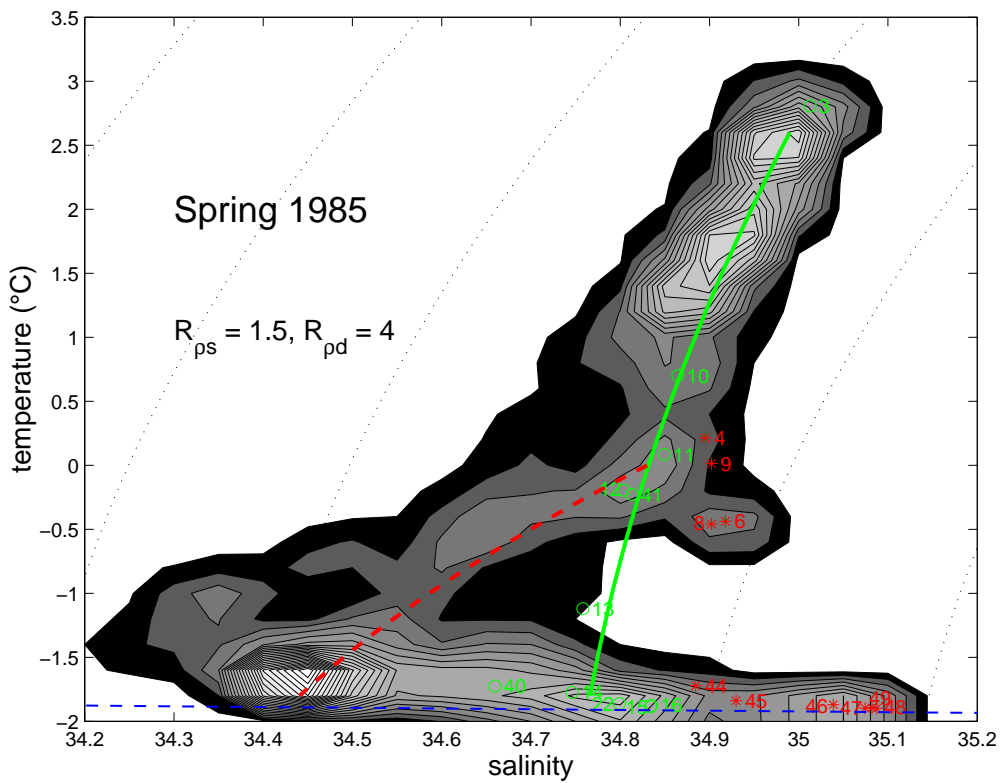
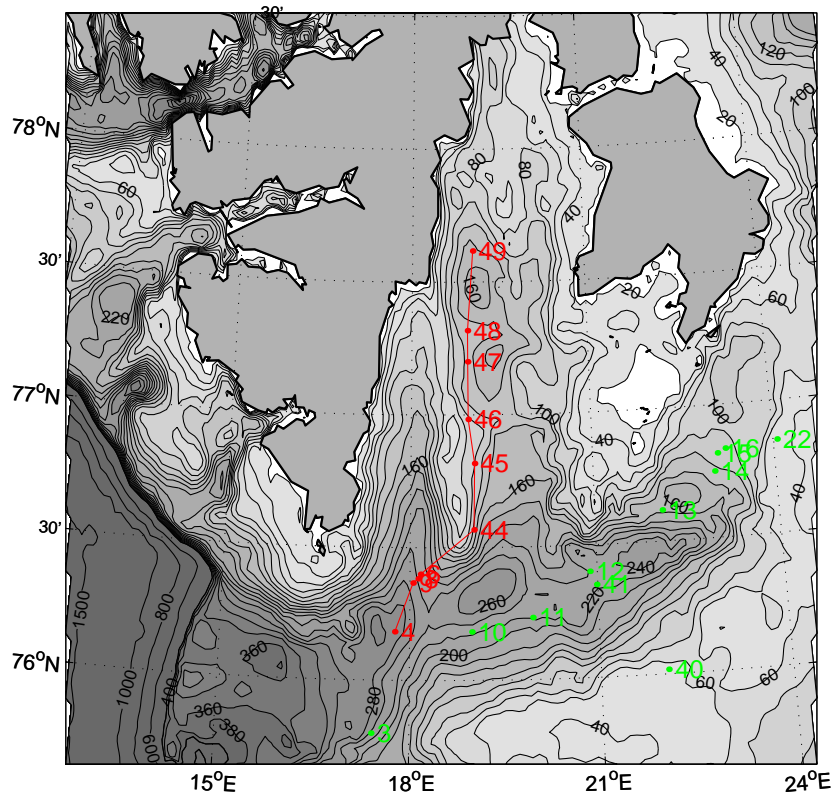
Andreas and Cash (1999, JGR)

Oceanic Double diffusion

Salt fingering
Thermohaline staircases



Rudels et al. (1999, J Mar. Syst.)



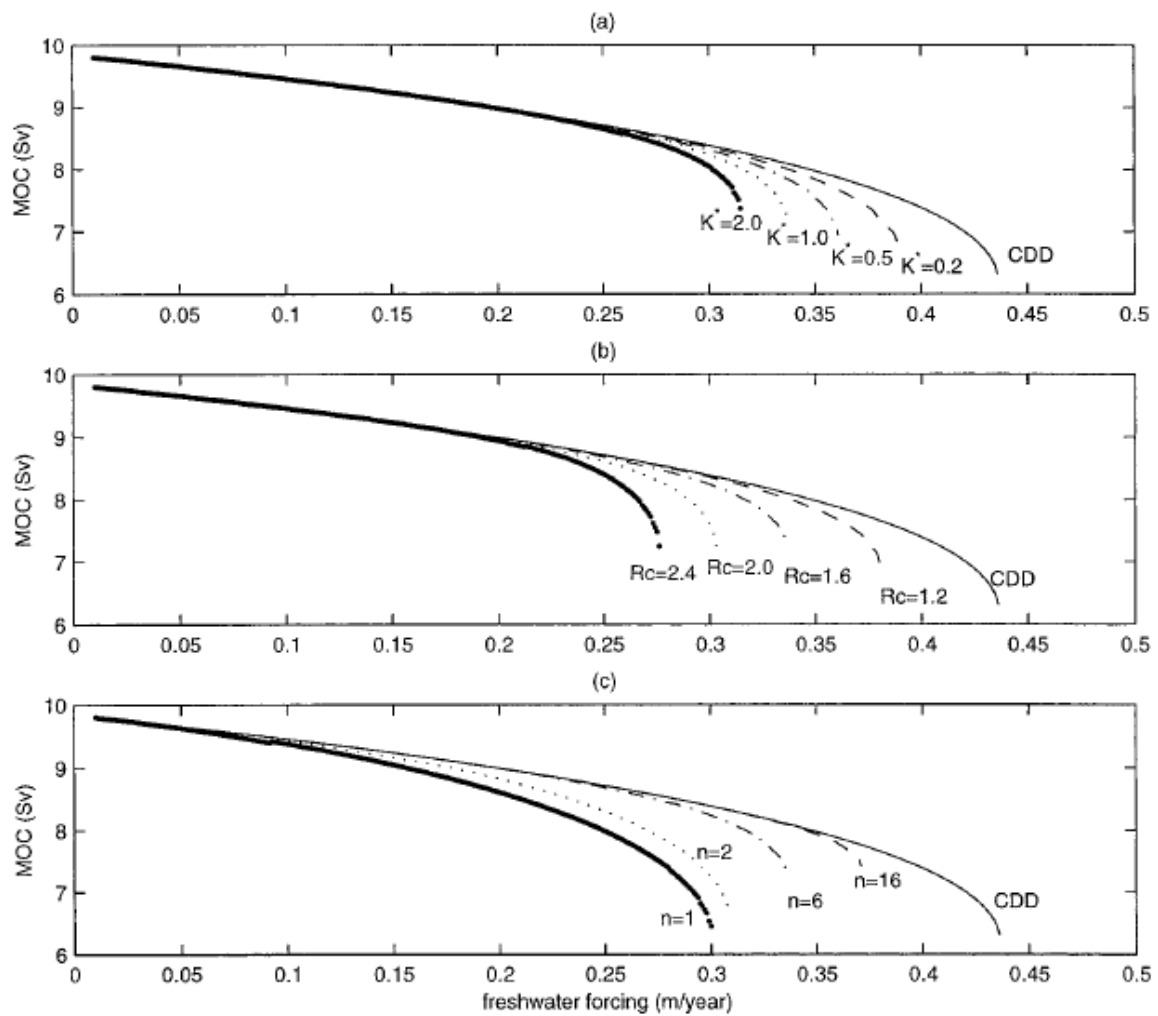
Ice melt

Salt fingering

Dense plume convection

Climate Models

Diapycnical mixing

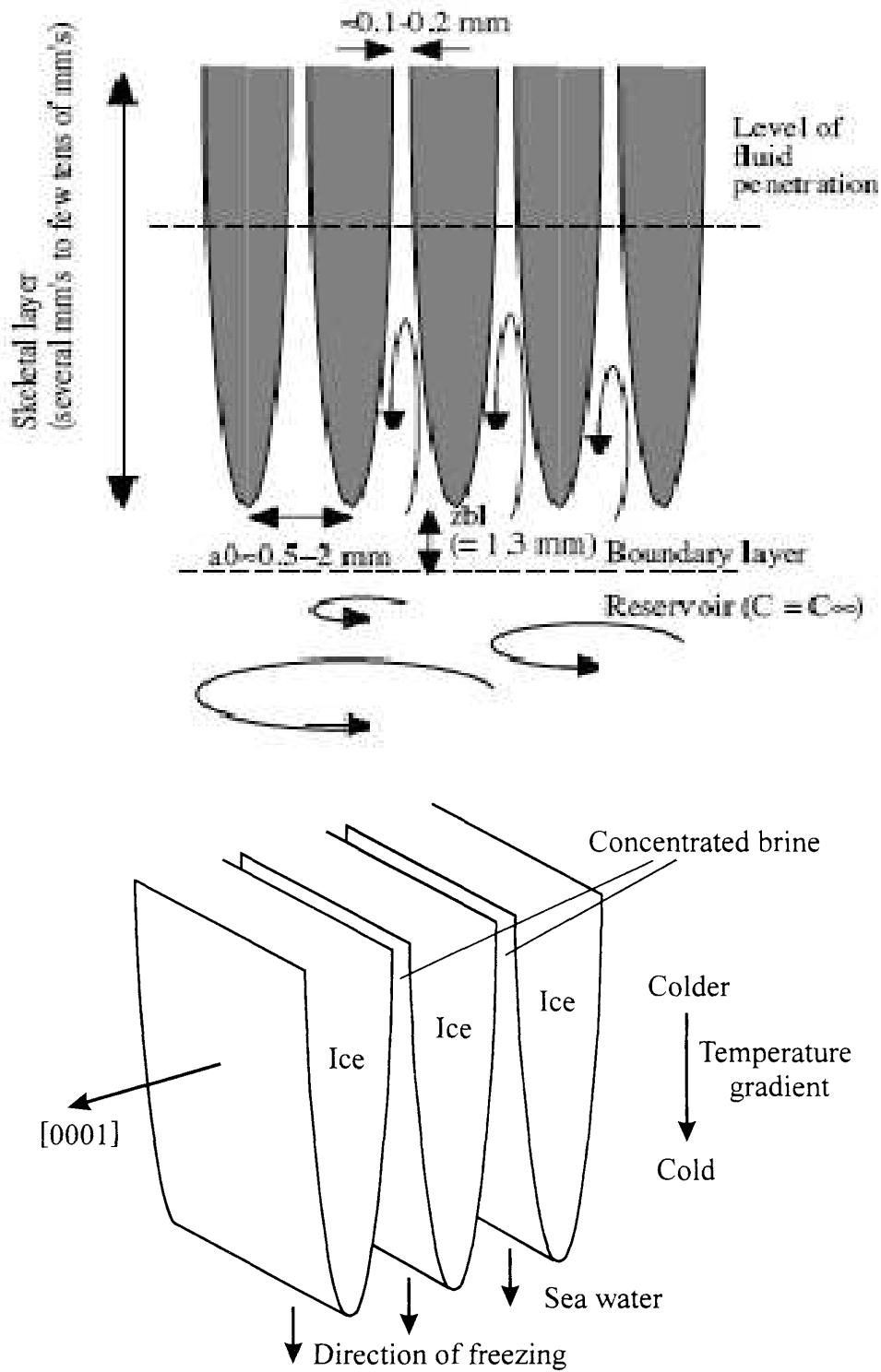


Zhang and Schmitt (1999, JPO)

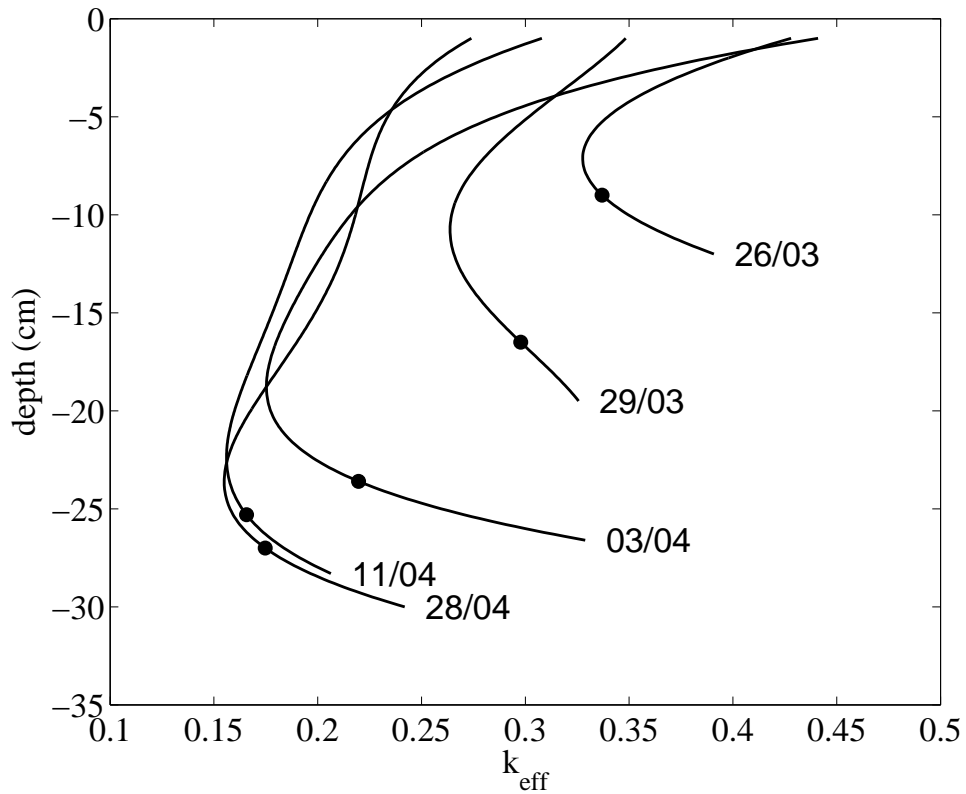
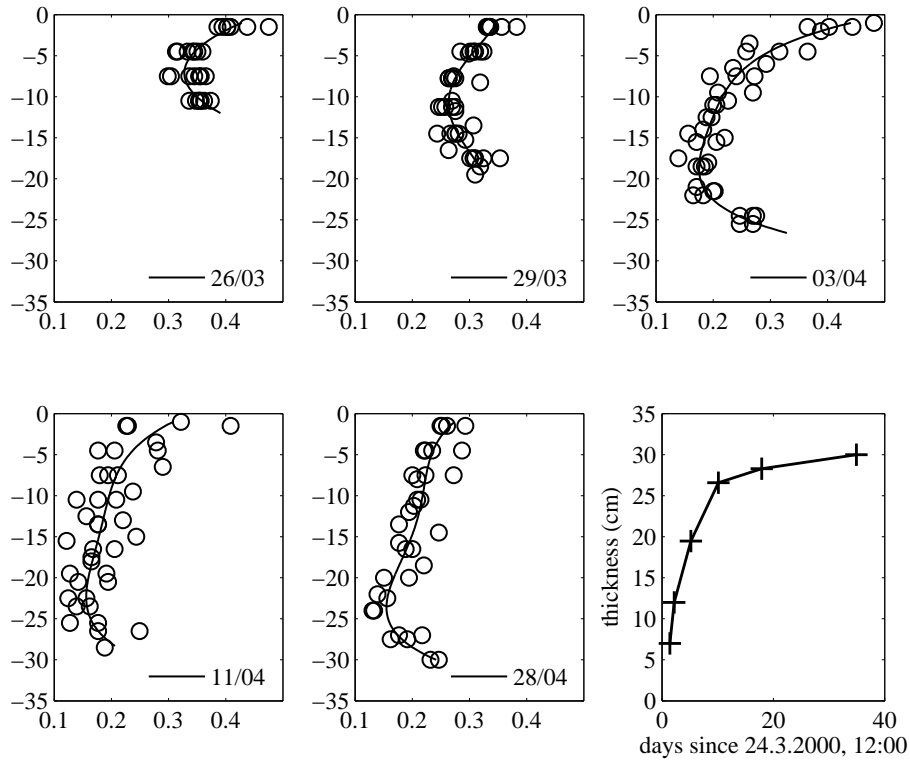
Sea ice freezing

Salt entrapment

Salt fluxes



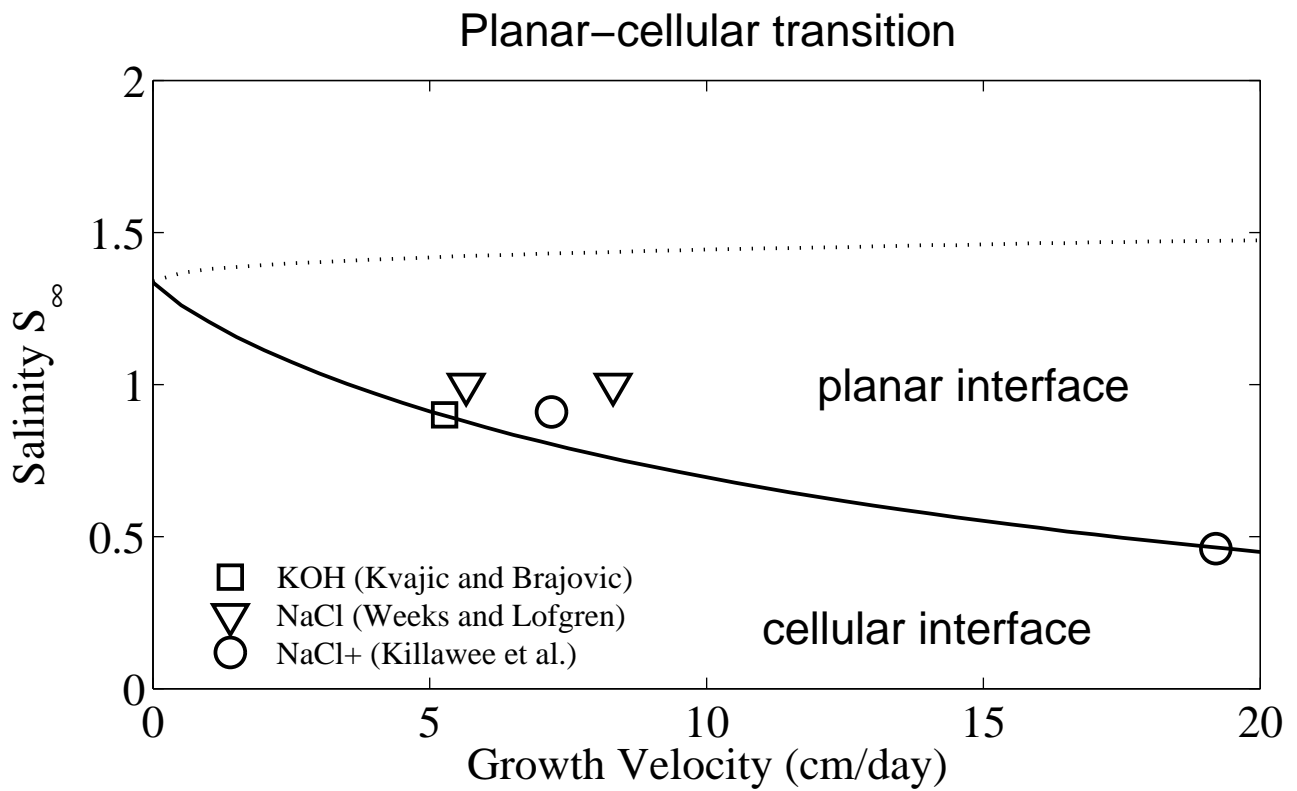
Salinity entrapment Salt fluxes



Adventfjorden observations 2000

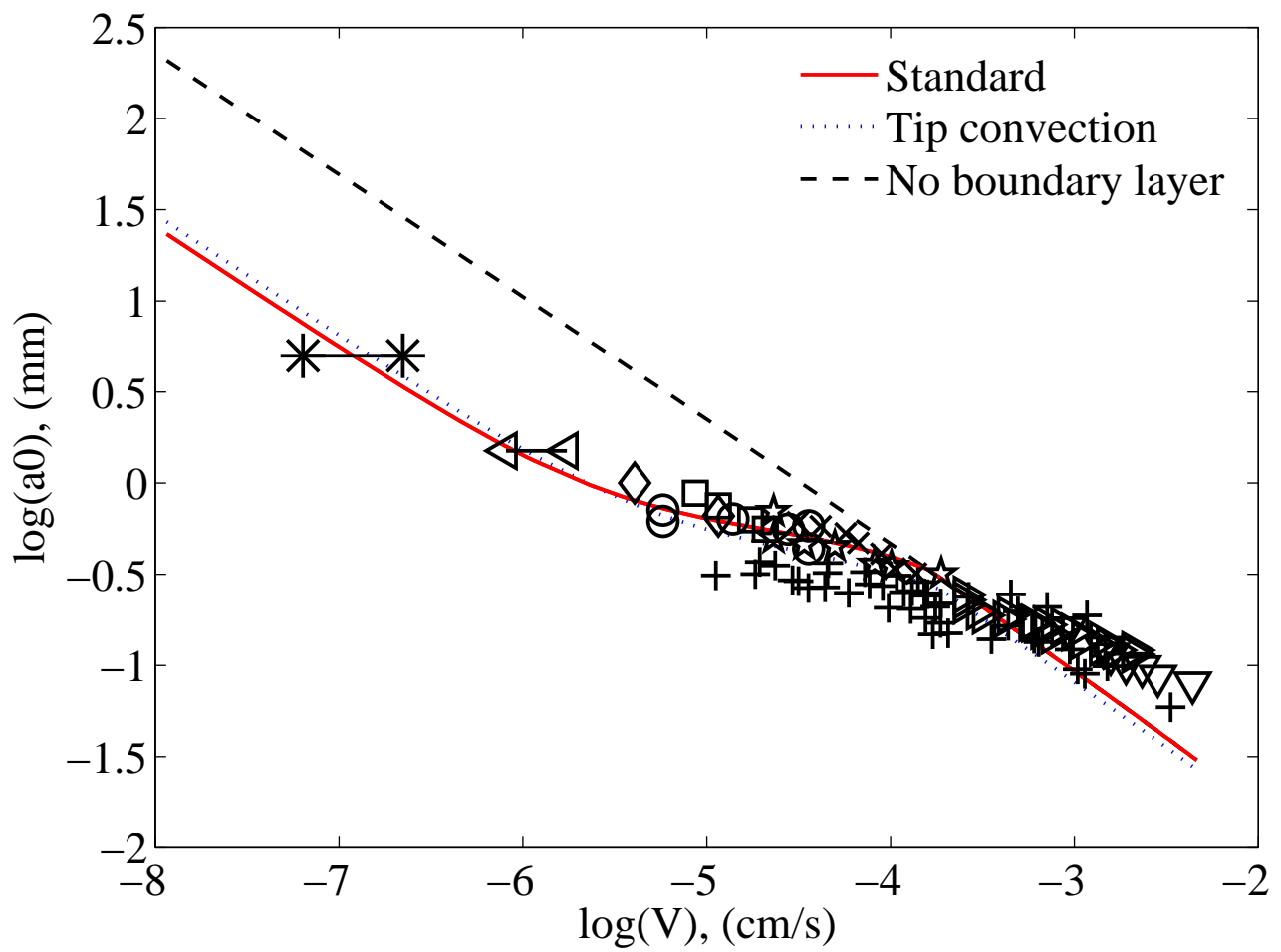
Planar-cellular transition

Fresh versus salt ice



Sea ice microstructure

Crystal growth



Sea ice microstructure

Sea ice salinity

Salt fluxes to the ocean

Ice shelves

